

AIMS Geosciences, 5 (1): 25–40.

DOI: 10.3934/geosci.2019.1.25 Received: 24 October 2018 Accepted: 07 March 2019

Published: 18 March 2019

http://www.aimspress.com/journal/geosciences

Research article

Soil organic carbon stock of different land uses of Mizoram, Northeast India

Alice Kenye¹, Uttam Kumar Sahoo¹,*, Soibam Lanabir Singh^{1,2} and Anudip Gogoi¹

- Department of Forestry, School of Earth Sciences and Natural Resource Management, Mizoram University, Aizawl-796004, Mizoram, India
- ² Dr Rajendra Prasad Central Agricultural University, Pusa, Samastipur-848125, Bihar, India
- * Correspondence: Email: uksahoo_2003@rediffmail.com, uttams64@gmail.com; Tel: +917005815370, +919436150944

Abstract: The study was conducted to assess soil organic carbon (SOC) concentration and stock under eight major land uses: shifting cultivation, wet rice cultivation, homegardens, forest (natural), grassland, bamboo plantation, oil palm plantation and teak plantation of Mizoram, Northeast India. Soil samples at different depths (0–15, 15–30 and 30–45 cm) were collected from each of the land uses under study to estimate SOC content in the laboratory. Forest recorded the highest mean SOC concentration with 2.74% at 0–45cm depth and lowest in the bamboo plantation (1.09%). Mean SOC stock for 0–45 cm soil depth ranged from 27.68 to 52.74 Mg C ha⁻¹ in grassland and forest respectively. Both SOC concentration and SOC stock decreased with increasing soil depth. Soil bulk density of fine soil (<2 mm) was significantly negatively correlated with SOC concentration and positive with SOC stock. SOC stock loss estimated following its conversion from forest was maximum with shifting cultivation (–5.74 Mg C ha⁻¹ yr⁻¹) followed by oil palm plantation (–2.29 Mg C ha⁻¹ yr⁻¹), bamboo plantation (–1.56 Mg C ha⁻¹ yr⁻¹) and the least in homegardens (–0.14 Mg C ha⁻¹ yr⁻¹). The study results indicate the importance of SOC stocks in different land uses which may help devise appropriate management practices to increase the soil carbon sequestration potential in the wake of mitigating climate change.

Keywords: land use types; bulk density; SOC stock; soil carbon sequestration

1. Introduction

Soil contributes largely to the global carbon cycle because it comprises of an active carbon pool [1]. In the terrestrial ecosystem, soil is considered to be the largest sink of organic carbon storing more than three times carbon compared to the amount stored in the atmosphere and 3.8 times more than the amount stored in biotic pool [2]. Therefore, the substantial sequestration of carbon in soils can provide a significant opportunity to mitigate global warming [3]. Enhancing the capture and storage of atmospheric CO₂ in different land use systems can be a successful approach to lower its concentration while also improving the quality of soil [4,5]. Soil profile in the top 1m stored 1500 Pg C soil organic carbon (SOC) globally, out of which Indian soil holds about 9 Pg C where the Himalayan zones account for about 33% of the total SOC reserves owing to thick forest vegetation [6]. SOC can either be increased or decreased depending on various factors such as soil type, climate, topography and soil management practices. However, SOC is greatly influenced by vegetation through organic matter input and therefore land use change is one of the most important factors which influences SOC stock build up. For example, it was reported that the conversion of farmland to apple orchard led to the decrease in the quality of soil owing to the reduced SOC stocks [7]. Soil carbon stock, following forest-pasture conversions, decreased to 51% in 20-30 years old pasture converted from wet tropical forest in Costa Rica, while SOC stock increased to 164% in a 33 years old leguminous pasture converted from native vegetation in Western Australia [8]. A meta-analysis reported that SOC stock decreased 13% and 42% when native forest converted to plantation and crop land respectively [9].

In natural ecosystem like forest and agroforestry, the soils are less disturbed due to less cultural operations and therefore may contain adequate nutrients and soil microorganisms when compared to agricultural lands [10–12]. Intensive management and cultural practices in agricultural lands increase the turnover rates of macro aggregates and lead to destabilization of the labile soil organic matter compounds [13]. Study reported from Northeast India showed the highest SOC stock in dense forest (140.4 Tg) and the least in shifting cultivation (10.7 Tg) with a total SOC stock (339.82 Tg), irrespective of the land use system for an area of 10.10 million ha, wherein forest soils contributed more than 50% with great implications for SOC sequestration in the region [14]. Studies from northern Bangladesh reported highest SOC concentration in agroforestry system (1.063%) and least in fallow land (0.249%) [15], whereas a similar study from homegardens in Aizawl, Mizoram reported SOC stock of 258.43 t C ha⁻¹ in 1 m soil depth [16]. Soil carbon sequestration proves to be a key indicator of soil health and crop efficiency [17,18], responsible for climate change mitigation and at the same time improving soil physical properties through moisture and nutrient retention [19]. However, the removal of biomass through deforestation and land use change can accelerate soil erosion resulting in significant loss of soil organic carbon from the surface soil [20,21]. The state of Mizoram reported a high percentage of forest cover (86.27% with respect to the total geographical area); however, forest cover has decreased considerably (by 531 km² from 2015 to 2017) due to shifting cultivation, biotic pressure, illegal felling, conversion of forest lands for developmental activities and agriculture expansion [22]. Despite the great potential of forest to sequester soil organic carbon, studies on SOC stock in forest and various land use conversions are limited in Mizoram [23–25]. Estimating SOC stock in various land uses has become very essential because it will aid policy makers to work out techniques for managing land use systems sustainably as well as preventing extreme loss of SOC. Hence, the present study was undertaken with objectives to estimate

SOC stock in different land uses and also to assess the relationship between SOC and land use types in Mizoram.

2. Materials and methods

2.1. Study site

This study was conducted in the whole of Mizoram which is located between 21°58' N to 24°35' N, and 91°15' E to 93°29' E encompassing a total area of 21,081 km² (Figure 1). The state is bounded internationally by Myanmar and Bangladesh on the southern part and domestically by Manipur, Assam and Tripura on the northern part. The climatic condition is mild with relatively cool summer 20 to 29 °C but becomes warmer with temperature exceeding 30 °C. In winter, the temperature varies between 7 to 22 °C. The winter season is short and summer long with heavy rainfall from the south-west monsoon with an average annual rainfall of 2450 mm. The monsoon period starts from May lasting till September with slight rain in the cold season. A summary of the site characteristics including climate, vegetation and soil of the eight land uses studied are shown in Tables 1 and 2. The age of the different land uses were determined with the help of the landholders and villagers.

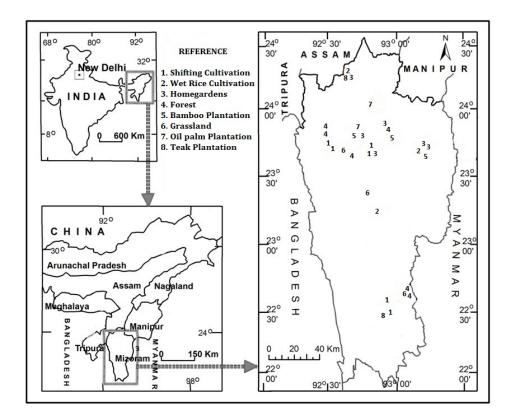


Figure 1. Map showing the location of different land uses considered for the study.

Table 1. Climate and vegetation characteristics of the 8 land uses studied.

	Land use							
Characteristics	SC	WRC	HG	Forest(Natural)	BP	GL	OPP	TP
Elevation	276–1658	39–1638	46–1682	562-2004	381–1532	792–1964	122–592	47–718
(m.a.s.l.)								
Slope	0–35	0–15	25–50	>50	25–50	25–50	25–50	15–25
(Min-Max %)								
Annual temp	17.3–24.76 °C	15.7–24.20 °C	15.7–23.83 °C	17.3–24.76 °C	15.7–24.20 °C	17.3–24.76 °C	18.85–24.20 ℃	19.7–24.76 °C
(Min-Max °C)								
Annual	2510-3155	2346-3067	2346-3155	2510-3155	2346-3155	2200-3067	2819-3067	2616-2819
rainfall (mm)								
Dominant	Musa	Oryza sativa	Parkia timoriana	Engelhardtia spicata	Melocanna	Quercus spp.	Elaeis	Tectona grandis
species	accuminata colla		Mangifera indica	Oroxylum indicum	baccifera		guineensis	
	Musa sylvestris		Artocarpus spp.	Helicia excelsia				
				Castanopsis tribuloidess				

SC—Shifting cultivation, WRC—Wet Rice Cultivation, HG—Homegardens, BP—Bamboo plantation, GL—Grassland, OPP—Oil palm plantation, TP—Teak plantation.

AIMS Geosciences Volume 5, Issue 1, 25–40.

Table 2. Soil characteristics of the 8 land uses studied.

Land use	Land use Characteristics						
types	Soil pH	Soil textural class	Soil colour	Major soil types	Parent material		
SC	Acidic	Sandy loam	Red and yellow loamy	Typic udorthents, Umbric-Dystrochrepts, Typic	Ferruginous sandstone, shale, alluvial		
				Dystrochrepts, Humic Hapludults, Typic	and colluvial materials		
				Hapludults			
WRC	Acidic	Sandy loam	Red and yellow loamy	Typic udorthents, Umbric-Dystrochrepts, Typic	Ferruginous sandstone, shale, alluvial		
				Hapludults	and colluvial materials		
HG	Acidic	Sandy loam	Red and yellow loamy	Typic udorthents, Umbric-Dystrochrepts,	Ferruginous sandstone, shale, alluvial		
				Humic Hapludults	and colluvial materials		
Forest	Acidic	Sandy loam	Red and yellow loamy	Typic udorthents, Umbric-Dystrochrepts, Typic	Ferruginous sandstone, shale, alluvial		
(Natural)				Dystrochrepts, Humic Hapludults, Typic	and colluvial materials		
				Hapludults			
BF	Acidic	Sandy loam	Red and yellow loamy	Typic udorthents, Umbric-Dystrochrepts,	Ferruginous sandstone, shale, alluvial		
				Humic Hapludults, Typic Dystrochrepts	and colluvial materials		
GL	Acidic	Sandy loam	Red and yellow loamy	Humic Hapludults, Typic Hapludults,	Ferruginous sandstone, shale, alluvial		
				Typic Dystrochrepts, Umbric-Dystrochrepts	and colluvial materials		
OPP	Acidic	Sandy loam	Red and yellow loamy	Typic udorthents, Umbric-Dystrochrepts	Ferruginous sandstone, shale, alluvial		
					and colluvial materials		
TP	Acidic	Sandy loam	Red and yellow loamy	Typic udorthents, Umbric-Dystrochrepts	Ferruginous sandstone, shale, alluvial		
					and colluvial materials		

SC—Shifting cultivation, WRC—Wet Rice Cultivation, HG—Homegardens, BP—Bamboo plantation, GL—Grassland, OPP—Oil palm plantation, TP—Teak plantation.

Volume 5, Issue 1, 25–40.

2.2. Soil Sampling and analysis

In each land use, five sample plots of $20 \times 20 \text{ m}^2$ were randomly selected, their locations and altitude recorded by a GPS. Within each sample plot, soils were collected from four corners and in the centre of the square plot at three depth classes: 0–15, 15–30 and 30–45 cm respectively. The five sub samples in each plot were mixed thoroughly and a composite sample was obtained for each depth class. A total of 120 samples (8 land use \times 5 plots \times 3 depths) were collected for SOC estimation. Similarly, a total of 120 samples (8 land use \times 5 plots \times 3 depths) for soil bulk density (BD) measurements were collected with the help of a soil corer of known volume. In the laboratory, the composite soil samples were mixed thoroughly, air-dried, crushed and passed through 2 mm sieve and replicates were made to analyse soil organic carbon content through Walkley and Black method [26]. For each depth, three replicates from each composite sample were analysed. Soil bulk density was determined by dry weight method by oven drying the soils at 80 °C for 24 hours and rocky fragments (>2 mm) were separated. Soil pH was measured using a pH meter and soil textural class was identified following ISSS soil mixture classification system.

2.3. Soil organic carbon stock estimation

Soil carbon stock for each site was estimated by multiplying with corresponding values of fine bulk density and SOC content. SOC stock was calculated following the formula given by Intergovernmental Panel on Climate Change (IPCC) [27].

C storage =
$$\sum_{\text{horizon}=1}^{\text{horizon}=n}$$
 (SOC × Bulk Density × Depth × (1 – frag) × 10)_{horizon} (1)

where, C storage—representative C stock for the soil of interest Mg C ha⁻¹. SOC—concentration of soil organic carbon in a given soil mass, g C kg⁻¹. Bulk Density—soil mass per sample volume, g cm⁻³. Depth—horizon depth or thickness of soil layer, m. Frag—% volume of coarse fragments (stone and gravel)/100, dimensionless.

2.4. SOC stock change estimation

SOC stock change (Mg C ha⁻¹) is estimated depending on the SOC stock changes between previous (C_{LU0}) and present (C_{LUn}) land use type [28]. The carbon stock of the previous land use type was set as the baseline for calculating the rate of change in SOC stock (Mg C ha⁻¹ yr⁻¹) after land use conversion. The following equation is used to calculate the rate of change (R_{stock}):

$$R_{stock} = \frac{C_{LU0} - C_{LUn}}{Age\ of\ C_{LUn}} \tag{2}$$

2.5. Statistical analysis

Descriptive statistics analysis was carried out using SPSS version 17.0. An analysis of variance (ANOVA) was done to evaluate if different land uses have significant SOC stock distribution and significant effect (P < 0.05) was determined with Tukey honest significant difference (HSD) post hoc multiple comparisons.

3. Results and discussion

Bulk density (BD) of fine soil (<2 mm) in different land use types across varying depths ranged between 0.40 to 0.71 g cm⁻³ (Figure 2). On an average, bulk density of fine soil in 0–45 cm soil profile was lowest in shifting cultivation (0.42 g cm⁻³) and the highest in forest (0.68 g cm⁻³). The higher soil bulk density in homegardens and wet rice cultivation as compared to other land uses maybe due to the cultivation practices such as tillage which cause soil compaction. Many studies have indicated that with the increase in soil depth, bulk density also tends to increase [29,30]. Conversely, our studies revealed that except in wet rice cultivation, the other land uses did not indicate any particular trend of bulk density increasing with increase in depth. It was reported that bulk density higher than 1.6 g/cm³ is unfavourable for plant growth as it can restrict the penetration of plant roots in clay loam soil [31]. In regard to this, the soil bulk density in all the land uses studied was found to be clearly below the critical value denoting that there is no extreme soil compaction.

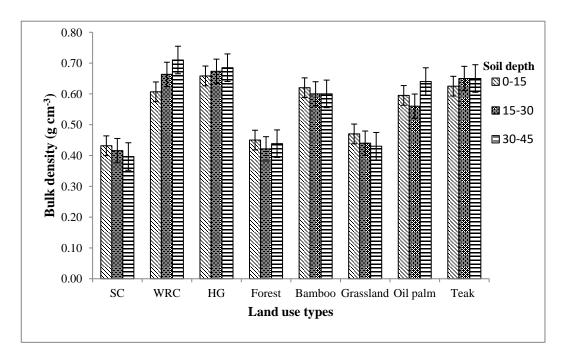


Figure 2. Depth wise soil bulk density in different land uses of Mizoram, Northeast India. (SC—shifting cultivation, WRC—wet rice cultivation, HG—homegarden). Error bars are standard errors of the means.

Amongst all the land uses studied, forest (natural) recorded the highest SOC concentration in all soil depths with 3.74, 2.70 and 1.79% in 0–15cm, 15–30 cm and 30–45 cm respectively. Bamboo forest recorded the least SOC concentration at 0–15 and 15–30 cm depth (1.28%, 1.10%) while grassland recorded the least (0.56%) at 30–45 cm depth (Table 3). SOC concentration decreased with increasing soil depth class in each of the land use types (Figure 3). On an average, forest soils have the highest SOC concentration compared to other land uses studied (Figure 4). The highest SOC concentration in the forest can be related to the presence of more vegetation generating more litter falls which are returned to the soils as organic matters. In the top surface soil (0–15 cm), maximum SOC stock was found in wet rice cultivation (26.36 Mg C ha⁻¹) followed by forest (24.50 Mg C ha⁻¹) and minimum in bamboo plantation (11.81 Mg C ha⁻¹) (Table 4). Many studies have shown that paddy soils have the potential to hold a great amount of recalcitrant/stable carbon [32,33]. In continuous

AIMS Geosciences Volume 5, Issue 1, 25–40.

wet rice cultivation, the built up of organic matter is very high due to their submerged conditions because in such a state, the decomposition of organic matter and SOC mineralization is very slow due to the anaerobic condition leading to higher net carbon storage [34]. Therefore, this may be one of the reasons why SOC stock in the upper 0-15 cm is higher in wet rice cultivation than the other land uses. In soil depth of 15-30 cm and 30-45 cm, the maximum SOC stock was recorded in homegarden and the least in grassland (Table 4). Grasslands have the ability to store a considerable amount of soil carbon in the upper stratum of the soil, however, the lower strata accumulates very less carbon which may be due to the shallow rooting of the grasses and absence of deep rooting trees. Highest SOC concentration and SOC stock in the top soil and decreasing with increase in soil depth from the study is in harmony with similar studies carried out by other researchers [11,23,35]. Overall amongst the different land uses, the mean SOC stock in 0-45 cm soil profile was highest in forest (52.74 Mg C ha⁻¹) followed by homegarden (50.85 Mg C ha⁻¹), wet rice cultivation (46.21 Mg C ha⁻¹), teak (44.66 Mg C ha⁻¹), oil palm (36.73 Mg C ha⁻¹), bamboo (29.83 Mg C ha⁻¹), shifting cultivation (27.87 Mg C ha⁻¹) and lowest in grassland (27.68 Mg C ha⁻¹) as presented in Figure 5. This is in disagreement with other studies where considerably high amount of SOC stock in grassland (75.76 Mg C ha⁻¹) than plantations (46.13 Mg C ha⁻¹) [36]; and higher value of SOC stock in grassland (95.54 Mg C ha⁻¹) than in agricultural land (75.70 Mg C ha⁻¹) [11] were reported. The significantly lower SOC stock in grassland reported in our study as compared to the other land uses might be because of the absence of deep rooted trees and fewer canopy covers. The potential for soils to store atmospheric carbon is primarily affected by the balance between the rate at which fresh photosynthetic material i.e. roots and exudates, is deposited and the time required by these carbon inputs to get broken down through heterotrophic respiration [37]. Moreover, as root tissue is more recalcitrant to degradation and mineralization than top soil litter, root derived carbon has long residence time [38]. Greater accumulation of organic matter in soils under tree canopies than open grassland can reduce leaching. Also, due to the absence of canopy covers, the soils in grasslands are directly exposed to the solar radiation thereby increasing the rate of mineralization. Similar findings have been reported by several other researchers [39,40]. Shifting cultivation recorded significantly lower (P < 0.05) SOC stock as compared to wet rice cultivation, homegardens and forest. Soils in shifting cultivation are usually much depleted due to reduction in biomass, reduced nutrients because of the shortened fallow periods and thus resulting in reduced soil organic carbon [41]. Another reason might be due to the steep slope condition of the state combined with heavy rainfall which leads to incidence of more soil erosion consequently leading to loss of SOC of the surface soils. A comparatively lesser SOC stock value of 29.83 Mg C ha⁻¹ in bamboo forest was observed from the study where other similar studies from Mizoram reported 46.04 Mg C ha⁻¹ [25]. This may be due to several reasons such as location, age of the bamboo stands and density of the bamboo stands. It was also reported by other studies that bamboo leaves releases allelopathic compounds during decomposition [42,43] which may reduce the growth of seedlings leading to low species richness which in turn reduces the input of litter thereby affecting the soil carbon stock. A correlation analysis of SOC concentration showed positive significant relationship with SOC stock, soil moisture content, clay and sand at P < 0.001level of significance. However, it correlated negatively with bulk density at P < 0.001 and silt at P < 0.05 level of significance respectively (Table 5). The positive relationship between SOC concentration and soil moisture content implies that SOC increases with increase in soil moisture content. This might be due to the microbial activity as soil moisture plays an important role in regulating the activity of soil microbes and determining the microbial population in forest floor [44]. Similar observations were reported by other studies too [45,46]. Furthermore, the significant positive relationship between SOC with clay and sand indicates the importance of fine soil particles for SOC storage for longer duration, especially clay minerals which protects against weathering and microbial degradation. Additionally, the relationship of SOC between different land use types is presented in Table 6. The significant positive relationship of SOC between different land use types may be due to similar land management practices such as in shifting cultivation and homegarden where the lands are regularly subjected to practices such as weeding and hoeing. Whereas, in case of shifting cultivation, grasslands and bamboo plantation, the lower input of litter due to less tree canopies might be one of the probable reasons. Forest exhibited a positive significant relationship only with teak plantation which may be attributed to the denser understory vegetation and less soil disturbances in both the land uses. Similarly, the significant negative relationship between oil palm with homegarden and bamboo; teak with grassland and oil palm can also be due to different management practices and types of inputs.

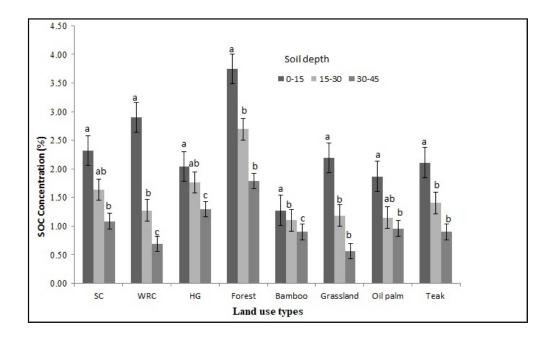


Figure 3. Mean SOC concentration in different soil depth per land use of Mizoram, Northeast India. Different letters indicate significant differences at P < 0.05. (SC—shifting cultivation, WRC—wet rice cultivation, HG—homegarden). Error bars are standard errors of the means.

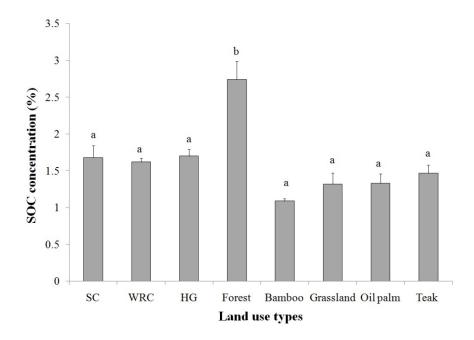


Figure 4. Mean SOC concentration (0–45 cm) in different land uses of Mizoram, Northeast India. Different letters indicate significant differences at P < 0.05. (SC–shifting cultivation, WRC–wet rice cultivation, HG–homegarden). Error bars are standard errors of the means.

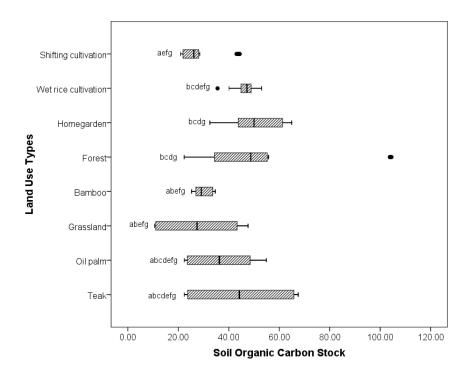


Figure 5. Soil organic carbon stock (Mg C ha⁻¹) in different land uses of Mizoram, Northeast India. Different letters indicate significant differences (P < 0.05) between the land uses. Error bars are standard errors of the means.

Table 3. Soil organic carbon (SOC, %) in different land uses of Mizoram, Northeast India.

	Soil Depth class (cm)					
Land use Types	0–15	15–30	30–45			
Shifting Cultivation	2.32 ± 0.24	1.64 ± 0.17	1.09 ± 0.11			
Wet Rice Cultivation	2.90 ± 0.05	1.28 ± 0.06	0.69 ± 0.05			
Homegarden	2.04 ± 0.10	1.76 ± 0.08	1.30 ± 0.08			
Forest (Natural)	3.74 ± 0.32	2.70 ± 0.30	1.79 ± 0.13			
Bamboo plantation	1.28 ± 0.04	1.10 ± 0.04	0.90 ± 0.03			
Grassland	2.20 ± 0.36	1.19 ± 0.12	0.56 ± 0.03			
Oil palm plantation	1.87 ± 0.25	1.15 ± 0.04	0.96 ± 0.13			
Teak plantation	2.11 ± 0.12	1.40 ± 0.19	0.90 ± 0.03			

 $[\]pm$ standard error of mean.

Table 4. Soil organic carbon stock (Mg C ha⁻¹) in different land uses of Mizoram, Northeast India.

		Soil Depth Class (cr	n)
Land use Types	0–15	15–30	30–45
Shifting Cultivation	13.14 ± 0.73	9.19 ± 0.63	5.55 ± 0.43
Wet Rice Cultivation	26.36 ± 0.86	12.61 ± 0.58	7.24 ± 0.47
Homegarden	19.95 ± 0.93	17.54 ± 0.61	13.36 ± 0.92
Forest (Natural)	24.50 ± 2.31	16.52 ± 1.91	11.71 ± 1.14
Bamboo plantation	11.81 ± 0.36	9.91 ± 0.34	8.11 ± 0.33
Grassland	16.09 ± 3.06	7.98 ± 1.12	3.60 ± 0.46
Oil palm plantation	17.29 ± 2.80	9.66 ± 0.42	9.76 ± 1.84
Teak plantation	20.57 ± 3.15	15.05 ± 3.37	9.03 ± 1.34

[±] standard error of mean.

Table 5. Pearson correlation (R) between SOC concentration and other parameters.

	SOC	SMC	рН	BD	Clay	Silt	Sand
SOC	1	0.375**	0.003	-0.324**	0.263**	-0.227*	0.336**
SMC		1	0.263**	-0.537**	0.100	-0.440**	0.511**
pН			1	-0.213*	-0.410**	-0.014	0.119
BD				1	-0.207*	0.304**	-0.250**
Clay					1	-0.689**	0.151
Silt						1	-0.089
Sand							1

^{*} Correlation is significant at P < 0.05 level (two-tailed).

^{**} Correlation is significant at P < 0.01 level (two-tailed).

Table 6. Relationship (Pearson's) of SOC (45 cm soil depth) between different land use types in Mizoram, Northeast India.

	SC	WRC	HG	Forest	Bamboo	Grassland	Oil palm	Teak
SC	1.00							
WRC	0.356	1.00						
HG	0.766**	0.016	1.00					
Forest	0.048	-0.493	0.328	1.00				
Bamboo	0.933**	0.259	0.862**	0.041	1.00			
Grassland	0.894**	0.125	0.885**	0.195	0.967**	1.00		
Oil palm	-0.664	0.524	-0.906**	-0.975	-0.891**	0.973**	1.00	
Teak	0.634	-0.491	0.899**	0.996**	0.904**	-0.977**	-0.971**	1.00

^{**} Correlation is significant at P < 0.01 level (two-tailed). SC—shifting cultivation, WRC—wet rice cultivation, HG—homegarden.

Meanwhile the estimated loss of SOC stock following conversion of forest to different land uses is presented in Figure 5. Result indicated a maximum loss of SOC stock when forest was converted to shifting cultivation (-5.74 Mg C ha⁻¹ yr⁻¹) followed by oil palm plantation (-2.29 Mg C ha⁻¹ yr⁻¹), bamboo plantation (-1.56 Mg C ha⁻¹ yr⁻¹) and the least in homegardens (-0.14 Mg C ha⁻¹ yr⁻¹). This indicates that emphasis should be given on proper management of shifting cultivation with respect to its intensity of practice and adoptability. The loss in SOC stock when forest is converted to other land use systems was also reported by other studies [8,9]. However, our research did not take into consideration of many other factors which are responsible for the gain or loss of SOC stock such as climate, altitude, soil types, physical, chemical, microbiological, biochemical properties of soil, etc. For instance, the plants community and productivity can be altered by the variations in climate along an altitudinal gradient which ultimately leads to the increase or decrease in the amount and the rate of soil organic matter (SOM) mineralization [47–49]. The decrease in temperature with increasing altitude has been proven to reduce the turnover rate of SOC in forest soils and thereby enhancing SOC stabilization and storage [50]. Furthermore, soil microbes play a huge role in the storage and stabilization of soil organic carbon as soil microbes can accelerate the rate of SOM mineralization. Studies reported that the higher microbial biomass in grassland comparing to arable land increased the mineralization of SOM [51]. Also, the increase or decrease in SOC stock may have been affected by indirect factors such as mycorrhizal colonization and soil aggregate size [52,53] Therefore, it is of utmost importance to consider these parameters while estimating SOC stock change in any land use system.

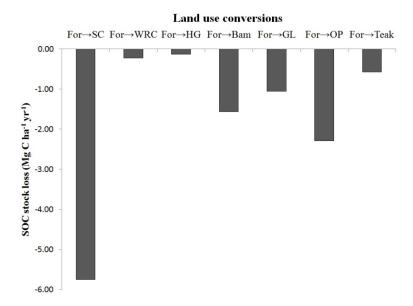


Figure 6. SOC stock loss following conversion of forest to different land use conversions in Mizoram, Northeast India. (For—Forest, SC—shifting cultivation, WRC—wet rice cultivation, HG—homegarden, Bam—bamboo, GL—grassland, OP—oil palm)

4. Conclusions

This study has shown that land use system is one of the major factors affecting soil organic carbon stock. It supports the existing knowledge of forest soils holding the maximum carbon stock. Our findings also indicated that a substantial amount of organic carbon can be stored in wet rice cultivation but shifting cultivation, which is a more dominant form of agriculture in Mizoram, resulted in a greater loss of SOC stocks due to their nature of practice and topography. Based on the results, SOC stocks in other land uses were low compared to forest and this indicate the presence of a good potential to sequester carbon in the soils of these land uses in the study area. Therefore, a detailed study of different land uses in Mizoram, and identification of its appropriate management practices aimed to increasing inputs and reduce soil organic carbon losses need to be conducted in the study area. Soil carbon sequestration will eventually minimize effects of climate change.

Acknowledgements

The authors would like to thank the Funding Agency: Department of Science and Technology, New Delhi under AICP- North-East CO₂ Sequestration Research Program of Ministry of Science and Technology, Government of India (DST/1S-STAC/CO₂-SR-227/14(G)-AICP-AFOLU-IV). The authors would also like to thank the editors and anonymous reviewers for their critical and valuable suggestions on this paper.

Conflict of interest

All authors declare no conflicts of interest in this paper

References

- 1. Prentice IC, Farquhar GD, Fasham MJR, et al. (2001) The carbon cycle and atmospheric carbon dioxide. In: Houghton JT, Ding Y, Griggs DJ, et al. (eds) *Climate Change 2001: The scientific basis, Contribution of Working Group I to the Third Assessment Report of the Intergovernmental Panel on Climate Change (IPCC)*, Cambridge University Press, 183–237.
- 2. Zomer R, Ustin S, Ives J (2003) Using satellite remote sensing for DEM extraction in complex mountainous terrain: Landscape analysis of the Makalu Barun National Park of eastern Nepal. *Int J Remote Sens* 23: 125–143.
- 3. Singh SK, Pandey CV, Sidhu GS (2011) Concentration and stock of carbon in the soils affected by land uses, soil types and climates in the western Himalaya, India. *Catena* 87: 78–89.
- 4. Lal R, Kimble JM, Follett RF, et al. (1998). The Potential of US Cropland to Sequester C and Mitigate the Greenhouse Effect, Ann Arbor Press, Chelsea, MI, 108.
- 5. Lal R, Follett RF, Kimble JM, et al. (1999) Management of U.S. cropland to sequester carbon in soil. *J Soil Water Conserv* 54: 374–381.
- 6. Bhattacharyya T, Pal DK, Chandran P, et al. (2008) Soil carbon storage capacity as a tool to prioritize area for carbon sequestration. *Curr Sci* 95: 482–484.
- 7. Shi Z, Li X, Zhang L, et al. (2015) Impacts of farmland conversion to apple (*Malus domestica*) orchard on soil organic carbon stocks and enzyme activities in a semiarid loess region. *J Plant Nutr Soil Sci* 178: 440–451.
- 8. Murty D, Miko UF, Kirshbaum F (2002) Does conversion of forest to agricultural land change soil carbon and nitrogen? A review of the literature. *Glob Chang Biol* 8:105–123.
- 9. Guo LB, Gifford RM (2002) Soil carbon stocks and land use change: A meta-analysis. *Glob Chang Biol* 8: 345–360.
- 10. Lemenih M, Karltun E, Olsson M (2005) Soil Organic Matter Dynamics after Deforestation along a Farm Field Chronosequence in Southern Highlands of Ethiopia. *Agroecosyst Environ*109: 9–19.
- 11. Sheikh I, Tiwari SC (2013) Sequestration of Soil Organic Carbon Pool under Different Land uses in Bilaspur District of Achanakmar, Chhattisgarh. *Int J Sci Res* 4: 1920–1924.
- 12. Gupta MK, Sharma SD, Kumar M (2014) Status of sequestered organic carbon in the soils under different land uses in southern region of Haryana. *Interl J Sci Environ Techn* 3: 811–826.
- 13. Six J, Elliott ET, Paustian K (1999) Aggregate and Soil Organic Matter Dynamics under Conventional and No-Tillage Systems. *Soil Sci Soc Am J* 63: 1350–1358.
- 14. Choudhury BU, Mohapatra KP, Das A, et al. (2013) Spatial variability in distribution of organic carbon stocks in the soils of North East India. *Curr Sci* 104: 604–614.
- 15. Iqbal MA, Hossen MS, Islam NM (2014) Soil organic carbon dynamics for different land uses and soil management practices in Mymensingh. *Proceedings of 5th International Conference on Environmental Aspects of Bangladesh*, 16–17.
- 16. Singh SL, Sahoo UK (2015) Soil carbon sequestration in homegardens of different age and size in Aizawl district of Mizoram, Northeast India. *NeBIO* 6: 12–17.
- 17. Yadav RL, Dwivedi BS, Prasad K, et al. (2000) Yield trends, and changes in soil organic-C and available NPK in a long-term rice-wheat system under integrated use of manures and fertilisers. *Field Crops Res* 68: 219–246.
- 18. Bolinder MA, Janzen HH, Gregorich EG, et al. (2007) An approach for estimating net primary productivity and annual carbon inputs to soil for common agricultural crops in Canada. *Agric Ecosyst Environ* 118: 29–42.

- 19. Leu A (2009) Applied organic systems, carbon farming and climate change. *Asian J Food Agro-Ind*, 307–317.
- 20. Sombroek WG, Nachtergaele FO, Hebel A (1993) Amount, dynamics and sequestering of carbon in tropical and subtropical soils. *Ambio* 22: 417–426.
- 21. Lal R (2002) Soil erosion and the global carbon budget. *Environ Int* 29: 437–450.
- 22. FSI (2017) India State of Forest Report. Forest Survey of India (Ministry of Environment and Forest and Climate Change), Dehradun, India, 248–253.
- 23. Singh SL, Sahoo UK, Gogoi A, et al. (2018) Effect of Land Use Changes on Carbon Stock Dynamics in Major Land Use Sectors of Mizoram, Northeast India. *J Environ Prot* 9: 1262–1285.
- 24. Singh SL, Sahoo UK (2018) Assessment of Biomass, Carbon stock and Carbon Sequestration Potential of Two Major Land Uses of Mizoram, India. *Inter J Ecol Environ Sci* 44: 293–306.
- 25. Devi AS, Singh KS, Lalramnghinglova H (2018) Aboveground biomass production of *Melocanna baccifera* and *Bambusa tulda* in a sub-tropical bamboo forest in Lengpui, North-East India. *Int Res J Environ Sci* 7: 23–28.
- 26. Walkley A, Black IA (1934) An examination of the Degtjareff method for determining soil organic matter, and a proposed modification of the chromic acid titration method. *Soil Sci* 37: 29–38.
- 27. IPCC (International Panel on Climate Change) (2003) LUCF sector good practice guidance. In: Penman J, Gytarsky M, Hiraishi T, et al., *IPCC Good practice guidance for LULUCF*. IPCC National Greenhouse Gas Inventories Programme, and Institute for Global Environmental Strategies (IGES), Hayama, Kanagawa, Japan, 3.1–3.312.
- 28. Deng L, Zhu GY, Tang ZS, et al. (2016) Global Patterns of the Effects of Land Use Changes on Soil Carbon Stocks. *Global Ecol Conserv* 5: 127–138.
- 29. Bessah E, Bala A, Agodzo SK, et al. (2016) Dynamics of soil organic carbon stock in the Guinea savanna and transition agro-ecology under different land-use systems in Ghana. *Cogent Geosci* 4: 1–11.
- 30. Mulat Y, Kibret K, Bedadi B, et al. (2018) Soil organic carbon stock under different land use types in Kersa Sub Watershed, Eastern Ethiopia. *Afr J Agric Res* 13:1248–1256.
- 31. Jones CA (1983) Effect of soil texture on critical bulk densities for root growth. *Soil Sci Soc Am J* 47: 1028–1121.
- 32. Liu QH, Shi XZ, Weindorf DC (2006) Soil organic carbon storage of paddy soils in China using the 1:1,000,000 soil database and their implications for C sequestration. *Global Biogeochem Cycles* 20.
- 33. Stern J, Wang Y, Gu B (2007) Distribution and turnover of carbon in natural and constructed wetlands in Florida Everglades. *Appl Geochem* 22: 1936–1948.
- 34. Sahrawat KL (2005) Fertility and organic matter in submerged rice soils. Curr Sci 88:735–739.
- 35. Brady NC, Weil RR (2008) *The Nature and Properties of Soils*, 14th Edition, Pearson Education, London.
- 36. Gupta MK, Sharma SD (2011) Sequestrated Carbon: Organic Carbon Pool in the Soils under Different Forest Covers and Land Uses in Garhwal Himalayan Region of India. *Int J Agric For* 1: 14–20.
- 37. Mathieu JA, Hatte C, Balesdent J, et al. (2015). Deep soil carbon dynamics are driven more by soil type than by climate: a worldwide meta-analysis of radiocarbon profile. *Glob Chang Biol* 21: 4278–4292.
- 38. Rasse DP, Rumpel C, Dignac MF (2005). Is soil carbon mostly root carbon? Mechanisms for a specific stabilisation. *Plant Soil* 269: 341–356.

- 39. Bernhard-Reversat F (1982) Biogeochemical cycle of nitrogen in a semi-arid savannah. *Oikos* 38: 321–332.
- 40. Isichel AO, Muoghalu JI (1992) The effect of tree canopy covers on soil fertility in a Nigerian Savannah. *J Trop Ecol* 8: 329–338.
- 41. Osman KS, Jashimuddin M, SirajulHaque SM, et al. (2013) Effect of shifting cultivation on soil physical and chemical properties in Bandarban hill district, Bangladesh. *J For Res* 24: 791–795.
- 42. Chou CH, Yang CM (1982) Allelopathic research of subtropical vegetation in Taiwan. II. Comparative exclusion of understory by *Phyllostachys edulis* and *Cryptomeria japonica*. *J Chem Ecol* 8:1489–1508.
- 43. Chan EH, Chiu CY (2015) Changes in soil microbial community structure and activity in a cedar plantation invaded by Moso bamboo. *Appl Soil Ecol* 91: 1–7.
- 44. Wagener SM, Schimel JP (1998) Stratification of ecological processes: a study of the birch forest Floor in the Alaskan taiga. *Oikos* 81: 63–74.
- 45. Das B, Bindi (2014) Physical and chemical analysis of soil collected from Jaismand. *Univers J Environ Res Technol* 4: 260–164.
- 46. Baishya J, Sharma S (2017) Analysis of Physico-Chemical Properties of Soil under Different Land Use System with Special Reference to Agro Ecosystem in Dimoria Development Block of Assam, India. *Int J Sci Res Educ* 5: 6526–6532.
- 47. Tate KR (1992) Assessment, based on a climosequence of soil in tussock grasslands, of soil carbon storage and release in response to global warming. *J Soil Sci* 43: 697–707.
- 48. Garten CT, Post WM, Hanson PJ, et al. (1999) Forest soil carbon inventories and dynamics along an elevation gradient in the southern Appalachian Mountains. *Biogeochem* 45: 115–145.
- 49. Quideau SA, Chadwick QA, Benesi A, et al. (2001) A direct link between forest vegetation type and soil organic matter composition. *Geoderma* 104: 41–60.
- 50. Trumbore SE, Vitousek PM, Amundson RR (1996) Rapid exchange between soil carbon and atmospheric carbon dioxide driven by temperature change. *Science* 272: 393–396.
- 51. Kara O, Bolat L (2008) The effect of different land uses on soil microbial biomass carbon and nitrogen in Bartin province. *Turk J Agric For* 32: 281–288.
- 52. Janzen H, Campbell CA, Brandt SA, et al. (1992) Light-fraction organic matter in soils from long-term crop rotations. *Soil Sci Soc Am J* 56: 1799–1806.
- 53. Manns HR, Maxwellr CD, Emery JN (2007) The effect of ground cover or initial organic carbon on soil fungi, aggregation, moisture and organic carbon in one season with oat (*Avena sativa*) plots. *Soil Tillage Res* 96: 83–94.



© 2019 the Author(s), licensee AIMS Press. This is an open access article distributed under the terms of the Creative Commons Attribution License (http://creativecommons.org/licenses/by/4.0)